

Groundwater Suitability for Irrigation: a Case Study from Debre Kidane Watershed, Eastern Tigray, Ethiopia

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ABSTRACT

The present paper tries to assess groundwater suitability for irrigation purpose in Debre Kidane Watershed (45.09 km²), northern Ethiopia. The goal was to evaluate the suitability of the groundwater for irrigation, examine the water types and to investigate possible long and short term impact on groundwater quality. Thirty six water samples each were collected in rainy and dry (irrigation) seasons from the active shallow hand dug wells. The well selection for detailed study and water sampling was done using stratified followed by random sampling technique. The number of wells selected for study is accounting 10% of the total available hand dug wells in the area. Parameters such as electrical conductivity (EC), pH, Total Dissolved Solids (TDS), temperature and other physical properties were recorded in the field. Major anions and cations (Ca²⁺, Mg²⁺, K⁺, Na⁺, CO₃²⁻, HCO₃⁻, Cl⁻, SO₄²⁻ and NO₃⁻) were analyzed in the laboratory. Chemical data suggests that the water type in the area though varies from Mg-HCO₃, Ca-HCO₃, Mg-Cl, Na-HCO₃ to K- HCO₃, predominantly it is mixed cation-HCO₃ type in the rainy season; and Ca-HCO₃ (Piper diagram) and Na-HCO₃ (Stiff diagram) in irrigation season. Dominance of Na in irrigation season and increase in Sodium Absorption Ratio (SAR) are primarily due to possible increase of Na absorption in the soil replacing Ca. According to the guidelines for irrigation water quality, the groundwater is suitable for irrigation with some minor exceptions. In rainy season, 89% of the samples fall under the water class “good” and 11% “permissible” whereas in irrigation season only 30% are classified as “good” and 70% under “permissible” class. Toxicity in terms of SAR and salinity though presently under control and there is no specific toxicity effect on vegetables and field crops, however, there are indications of their increase in due course of time if proper measures are not taken.

Keywords: Groundwater, Irrigation, Salinity, SAR, Toxicity, Water quality, Hand dug well, Ethiopia.

1. INTRODUCTION

Water occurring either as surface or subsurface water is the mankind's most vital and versatile natural resources. Surface water exists in rivers, lakes, ponds, and oceans. Surface water mainly rainwater by percolating through soil and weathered rock horizons occupies the subsurface permeable layers and develops into a groundwater. Due to its interaction for longer period of

time with soil, weathered and fresh rock materials which in turn vary in composition, groundwater compared to surface water provides a wide range of variation in its composition. Thus, the utility of groundwater depends not only on its abundance but also on its quality. Groundwater utilization for irrigation and domestic purposes has been at a maximum compared others, like livestock and, industrial use etc., particularly due to food security problems and global climatic changes. Thus groundwater has become an important source for irrigation and it has become an integral part of the irrigation strategy “to overcome food scarcity” in many developing countries including in African continent. Such efforts have helped to expand irrigation with time particularly from 1950’s the expansion has been rapidly increasing (Rosegrant et al., 1999). Currently, it accounts for about 72% of global water withdrawals and about 90% in the case of developing countries.

According to the Ministry of Water Resources (2002), Ethiopia is endowed with huge natural water resources, which include 122 billion m³ annual surface runoff and 2.9 billion m³ of groundwater. However, the country’s water resource has contributed little to the country’s socio-economic development, because major part of the surface runoff is not utilized and at the same utilization of groundwater for irrigation is still in infantile stage. In the case of the Tigray region the main economic means is rain fed agriculture. The region is characterized by undulating topography and experiences arid to semi-arid conditions with highly erratic and unreliable rainfall. Thus the region is not in a position to cover the annual food requirement of its people. To alleviate the challenges of food insecurity in the country promotion of irrigated agriculture was given priority (Mekuria, 2003). Thus, hand dug well construction has become one of the activities both by the individuals as well as the government on a sustainable basis. Implementation of this technology does not need high investment and skilled work force instead a household can easily adopt and practice it. However, the issue of sustainability in terms of quality and quantity which changes with time and type of practices in the field, demands attention.

The changing conditions have provided impetus to develop irrigation using groundwater. Debre Kidane Watershed in Tigray is one such area where about 360 shallow hand dug wells were constructed primarily to overcome the moisture scarcity. The households benefited from these dug wells and started cultivating and producing different high value crops even two times a year. Availability of water by itself is not a guaranty for sustainable development, but, its fitness to

specific purpose like irrigation, domestic and industries is important. Since economic advantage alone cannot sustain the practice, the issues like ecological and social are to be taken in to account in the development and management of the irrigation schemes. Quality of water should be made part of such studies. Ultimately the knowledge of irrigation water quality is critical to understand what management changes are necessary for long-term and short term productivity particularly for crops that are sensitive to changes in quality (Bohn et al., 1985; FAO, 1985; Brady, 2002).

The present paper examines the quality of groundwater used for irrigation in Debre Kidane watershed (Fig.1) and tries to a) determines suitability of groundwater for irrigation purpose; b) investigate the water types; c) record water salinity, sodicity if any; d) determine the possible sources for such variations; and, e) examine possible future trends in the toxicity.

1.1 Description of the Study Area

Debre Kidane watershed is located in the eastern Tigray, North Ethiopia. Geographically, it is found bounded between 545000 – 553000m E and 1535000 – 1544000m N, and has an aerial coverage of 45.09 km² (Fig. 1) with a mean altitude of 2200 meters above sea level. The mean annual rainfall in the area is about 524.08 mm (Nata, 2003). Monthly rainfall is concentrated mostly in the mid of June to the mid of September. The drainage is well developed and originates from the elevated areas in the northeast and flow towards south west (Fig. 1). Northeastern elevated areas form the main recharge source for groundwater in the area. The mean annual temperature is 18.1°C and the yearly average maximum and minimum temperatures are 25.1°C and 10.8°C, respectively. The annual range of temperature is 3.7°C (Nata, 2003). The dominant crops grown in the area are barley (*Hordeum vulgare*), wheat (*Triticum aestivum*), teff (*Eragrostis tef*), maize (*Zea mays*), finger millet (*Pennisetum americanum*), beans (*Phaseolus vulgaris*) and peas (*Pisum sativum*).

Soil texture observed in the field are sandy, light clay, silty, loamy sand, and loam. Dominant soil texture being silt loam (Nata et al., 2008). Major soil of the watershed from along the inlet and outlet is Vertisols, Cambisol, Leptosol, and Alluvisol.

In the study area the types of aquifers are localized, confined and unconfined types. Confined aquifers are located in the discharge areas of the basins and the unconfined types are found in the recharge areas. Out of the observed hand dug wells, 72.5% are located in unconfined aquifers and the remaining are in the confined aquifers (Nata et al., 2008).

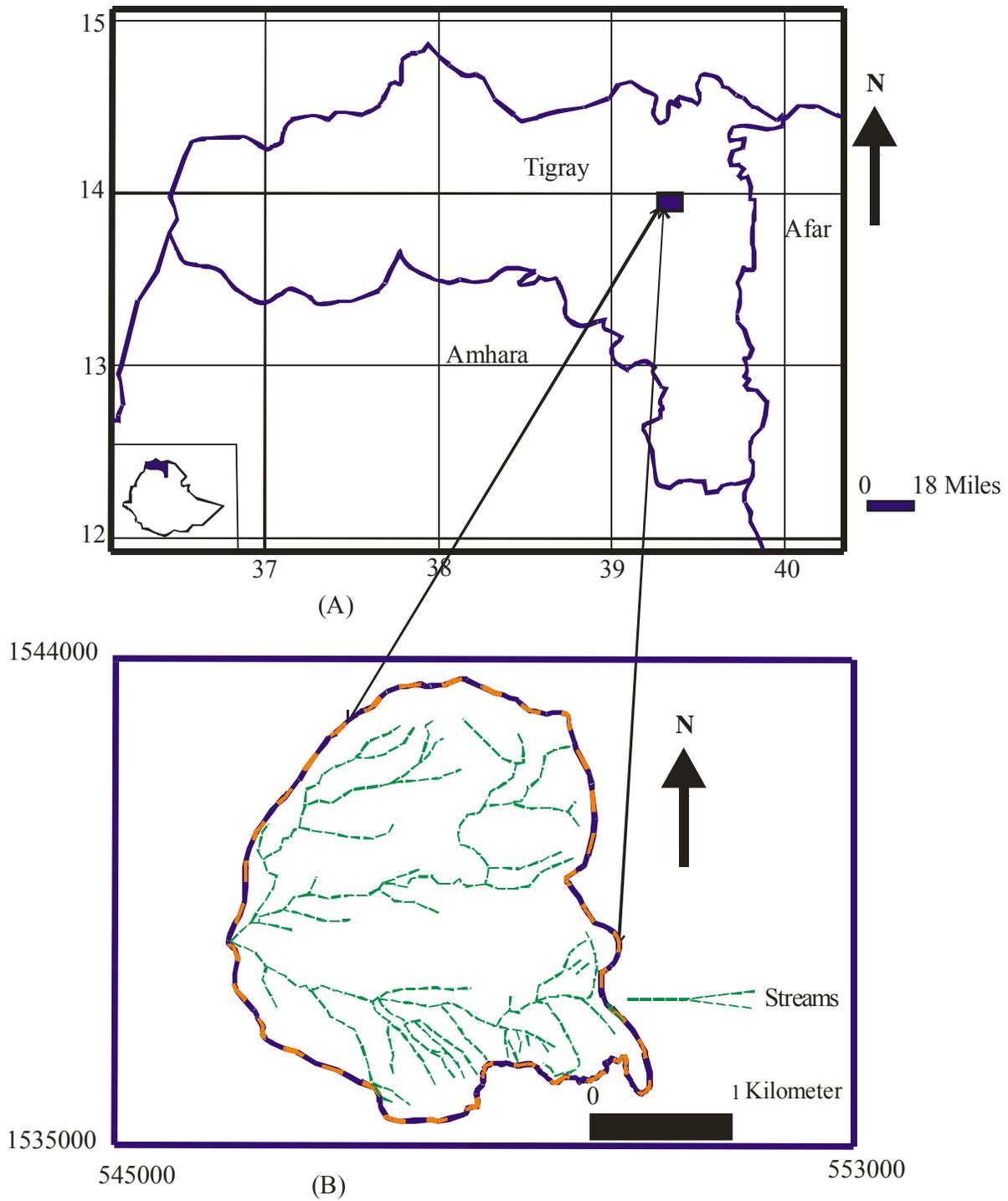


Figure 1. Location and drainage map of the watershed (Nata et al., 2007).
(Source for A: Disaster Prevention and Preparedness Agency (DPPA), 2006)

2. GEOLOGY AND AQUIFER CONDITION OF THE STUDY AREA

The area consists of three types of rock units, the low-grade metamorphosed Proterozoic basement rocks and overlying sedimentary rocks comprising sandstone and tillite of Paleozoic age and recent alluvial deposits (Fig. 2) (Nata et. al., 2007). A brief account of the rock types and the aquifer characteristics is given below.

2.1 Geology

2.1.1 Basement rocks

These are the dominating rock type in the area and their outcrops are found in northern, northeastern, eastern, southeastern, central, southwestern, and northwestern parts of the study area. They occupy both low in the southern parts and high elevated areas in the northern parts and show flat to steep slopes. They cover 22.8% of the area (Nata et. al., 2007).

The basement rocks are composed of metavolcanics and metasediments. The most conspicuous foliation trend is north-south with deviations to northeast and northwest. Apart from foliations, fractures are also present and are irregularly spaced in the rock. In some places they show light to dark grey, light to dark brown and reddish color. Dominant minerals present in these low grade greenschist facies rocks include: mafic minerals, chlorite, quartz, feldspars and opaque in metavolcanic rocks and quartz, chlorite, feldspar, muscovite, pyrite, limonite in metasediments (Nata et. al., 2007).

2.1.2 Sandstone

This rock unit accounts for about 8.4% of the total study area (Nata et. al., 2007). The outcrops of these rocks are found in the southern and southeastern parts of the study area. They occupy higher elevations in the topography, show steep slopes, cliff forming, and overlie the basement rocks. The rock out crops is very limited as they are covered by alluvium. However, as it was observed from the geological logs of the hand dug wells, it extends from southeast to the central, southwestern, western and northwestern parts of the study area. Its maximum thickness which is observed at the southeastern periphery of the watershed is of about 160 meters (Nata et. al., 2007).

It is composed of white, coarse grained, cross-bedded, calcareous sandstone containing lenses of siltstone, grit and polymict conglomerate with sub-rounded to well-rounded pebbles, cobbles and boulders. Clasts of granite, gneisses are scattered erratically and are seen in many places. The lower part of the sandstone is white in color with medium grained angular grains and at places

strongly cemented by clay minerals. Upper part is the glacial unit, shows features of a glacio-fluvial deposit, angular, poorly sorted grains of quartz and lenses of conglomerate. The rock is porous, highly weathered and show presence of fractures, joints and cross bedding. Among structures, joints are prominent and show presence of two sets of joints in sandstone which lies N15°E and N25°W (Nata et al., 2007).

The tillite unit is not found as an outcrop in the study area. However, as it was observed from the lithologs of the hand dug wells, this rock is found underlying the alluvium in the northern, northeastern, eastern, southeastern, southern southwestern parts of the watershed (Nata et al., 2007). It is characterized by dark grey color, poorly sorted texture, the presence of big boulder and its unstratified nature. On the basis of the logs of the hand dug wells, its maximum thickness is about 4.5m. The rock is dominated by quartz with limited amount of feldspar and varying amounts of clay, silt and ferruginous and calcareous cementing material (Nata et al., 2007).

2.1.3 Alluvial Deposits

In the watershed alluvial deposits are found covering wide area both in the highlands and the lowlands. They are also found as thin strips along the margins of the major rivers and their tributaries. The alluvium is accounting for about 68.8% of the area. The relative abundances and stratigraphic relations of the sediments, however, are generally not uniform through out the watershed. Toward the mountains front that is NW and E and topographically high areas in the central parts of the study area, where river gradients are high due to steep slopes, the alluvial sediments, in general, are dominated by sub-angular to sub-rounded coarse grained quartz. In the western parts of the watershed, where the gradient of the rivers decreases down slope, the alluvial deposit is characterized by the presence of medium to fine grained sand with variable content of silt and clay. The maximum thickness of the alluvium is about 3.50 m.

2.2 Hydrogeology

2.2.1 Aquifer Types and Characteristics

Different rock types and unconsolidated sediments which host water and act as aquifers have been classified into two based on the type of porosity and permeability and their extent.

1. Localized aquifers with intergranular porosity and permeability (unconsolidated sediments: alluvial sediments along the margins of the major river and its tributaries); and,

2. Less extensive aquifers with intergranular and fractured porosity and permeability (consolidated sediments and basement rocks: sandstone, tillite and basement rocks).

A brief account of the hydrogeological characteristics of the different rocks and unconsolidated sediments of the watershed is given below with a particular reference to their water storage and transmission capacities (Fig. 3) (Nata et al., 2008).

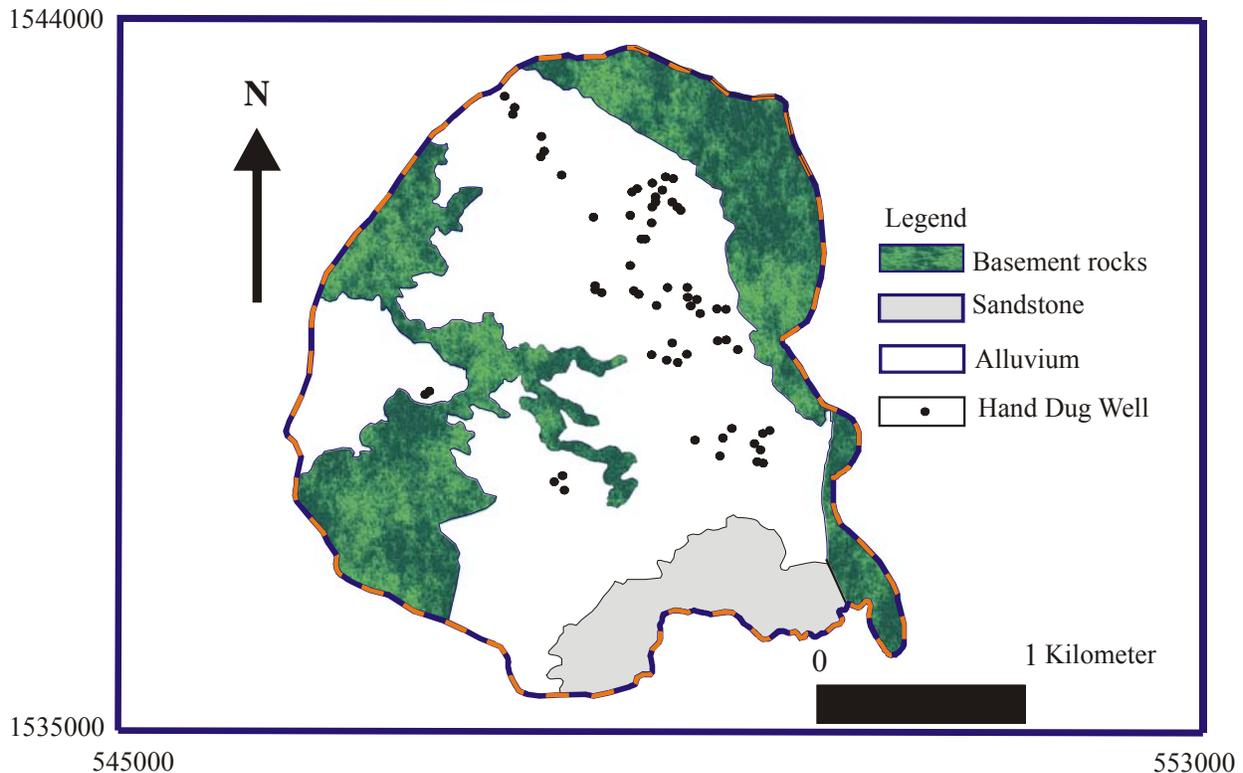


Figure 2. Geological and Hand dug well location map, Debre Kidane, eastern Tigray (Nata et al., 2007).

2.2.1.1 Basement rocks

These rocks are generally impervious, nevertheless local pockets of groundwater reservoir occurs in the weathered layer and fracture zones. Geological and hydrogeological logs of nine hand dug wells in the western and central parts of the watershed indicate that the weathered and fractured zones are the main sources of groundwater supply in the basement rocks with discharges of hand dug wells range from 0.4 to 1.5 l/s.

2.2.1.2 Sandstone

The rock is characterized by highly weathered and fractured nature. These weathered and fractured zones as a whole reaches up to a maximum depth of about 4 m. However, in this rock secondary porosity is significant due to fracturing. Fracturing has increased the void space as well as its capacity for water transmission and enhanced its usefulness to the water supply. However, at places it is reduced due to high degree of cementation and in turn permeability and productivity. This rock is considered to host a highly potential aquifer in the flat areas (northern and northeastern parts of the watershed) compared to the steeply to gentle dipping rocks which act as a conduit rather than being an aquifer in the southern and southeastern parts of the area.

In the case of tillite presence of considerable amounts of silt and clay as minerals and as cement reduces the intergranular permeabilities drastically and makes the rock impervious and hamper water movement. At places, however, the fractures provide secondary porosity and permeability and hence make the rock permeable for groundwater flow and storage,

2.2.1.3 Alluvial Deposits

These recent deposits cover major parts of the area overlying both the basement and sedimentary rocks. They vary in thickness as well as in composition in the plateau areas and along the margins of the rivers and tributaries. The alluvial deposits mainly comprise of clay, silt, sand and gravel sized particles in different proportion. These deposits have an average thickness 3.50 m. In the western part of the study area, these deposits can be considered as potential water bearing formation due to their primary porosity and location whereas due to their location in the southern parts the alluvium act as a conduit rather than being an aquifer (Fig. 3).

3. METHODOLOGY

Initially inventory of all hand dug wells that are functional for irrigation purpose was carried out. During inventory in situ measurement of electrical conductivity, temperature of the groundwater and air temperature for each well were carried out. Since the electrical conductivity values were measured *in situ* at a temperature different from the standard 25 °C, an adjustment of the electrical conductivity values of the water was made by multiplying the respective measured electrical conductivity value by the factor corresponding to the temperature at which the measurement was made.

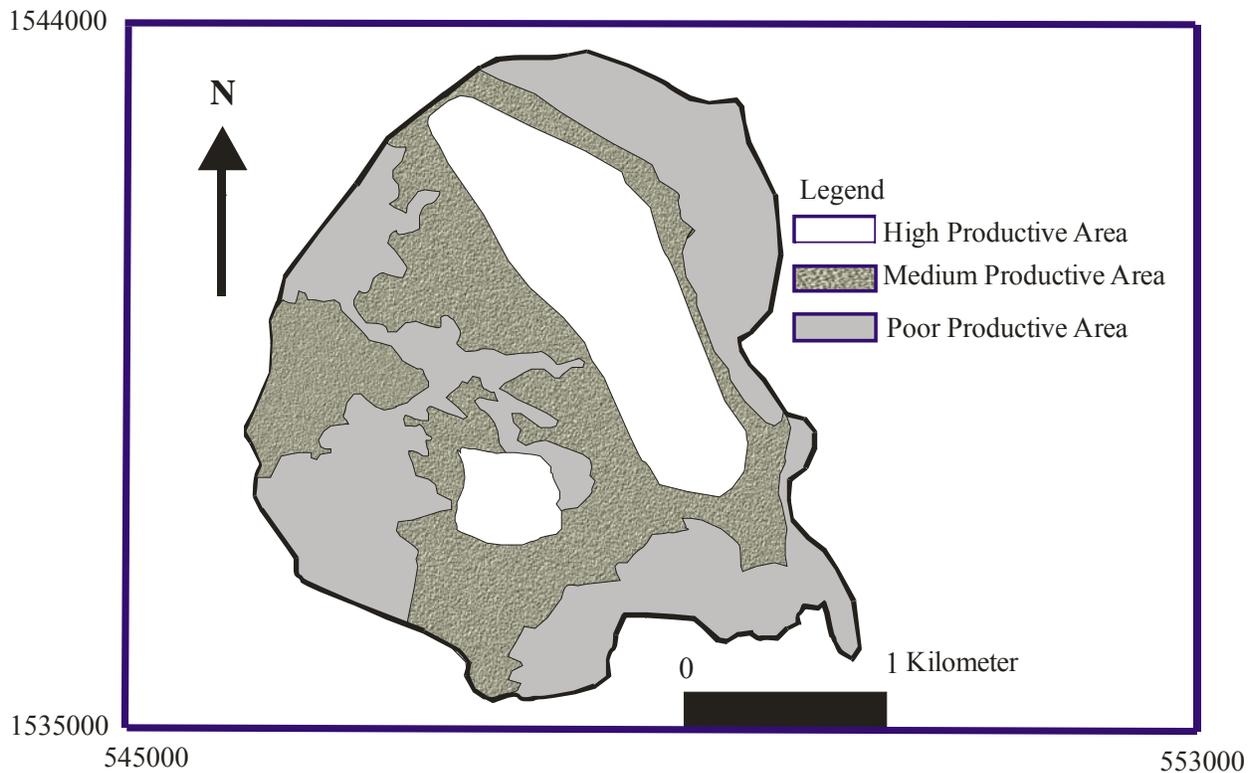


Figure 3. Groundwater potential in the study area (After Nata et al., 2008).

To determine the number of water samples for chemical analyses stratified and random sampling techniques were utilized. The *in situ* measured and corrected electrical conductivity values of the groundwater were grouped into different water classes based on Quality Classification of Water for Irrigation (Wilcox, 1955). Then after, 36 water samples were selected randomly from the different water classes for chemical analyses. Since majority of the wells in the area are present in the alluvium in the river valley, the wells selected for sampling are also restricted to this part of the area. Thirty six water samples each were collected from hand dug wells in both rainy season and irrigation (dry) season. Apart from temperature, other physical parameters like turbidity, colure, taste and odor were also measured at each site. For the purpose of chemical analysis the water samples were collected in one liter plastic bottles after thorough wash. All the water sample locations are shown in the figure 2.

The water samples were analyzed in the Geochemistry Laboratory of Applied Geology Department, Mekelle University, Mekelle. The samples were analyzed for calcium (Ca^{2+}), magnesium (Mg^{2+}), potassium (K^+), sodium (Na^+), bicarbonate (HCO_3^-), chloride (Cl^-), sulfate

(SO_4^{2-}), and nitrate (NO_3^-). Besides, pH and electrical conductivity in $\mu\text{S}/\text{cm}$ at 25°C were also measured.

The cations were analyzed using Atomic Absorption Spectrophotometer and the anions using UV-Spectrophotometer (Varian Spectra, 50B). The data are presented in the form of graphs. Piper, Stiff and Box and Whisker diagrams were used for representing and comparing water quality analyses

In this study FAO (1985 and 1989) guidelines are used to evaluate the suitability of the groundwater of the watershed for irrigation. The guidelines were proposed for evaluating the potential of an irrigation water to create soil or crop problems.

Various thematic maps such as location, digital terrain model of the watershed, drainage, geological, and hydrogeological were produced by using ArcView 3.3 and CorelDRAW 12 software's.

4. RESULTS

4.1 Electrical Conductivity, pH and Total Dissolved Solid (TDS)

EC values were recorded in 36 functional hand dug wells indicate that electrical conductivity values range from $0.05\text{-}2.0\text{ dSm}^{-1}$ with a mean of 0.945 dSm^{-1} . The values measured for 36 groundwater samples vary from 0.27 dSm^{-1} to 0.96 dSm^{-1} with an average of 0.53 dSm^{-1} at 25°C for rainy season. The same during irrigation season range from 0.399 to 1.696 dSm^{-1} at 25°C with a mean of 0.877 dSm^{-1} at 25°C . The EC values in groundwater are much higher during irrigation season up to 65% compared to the same of the rainy season (Fig. 4).

Hydrogen ion concentrations (pH) range from 7.70 to 8.60 during rainy season and 7.22 to 8.39 during irrigation season. The average pH value in rainy and irrigation season is 8.12 and 8.01, respectively. Total dissolved solid (TDS) values in both the seasons for groundwater range from 274.34 to 1092.2 mg/l. In general, TDS values for groundwater in both rainy and irrigation seasons are less than 1000 mg/l.

4.2 Cations

Among cations, calcium is the most abundant element and varies from 8 to 60 mg/l during rainy season with a mean concentration of 35.03 mg/l. During irrigation season, the most abundant cation is Ca^{2+} followed by Na^+ , K^+ and Mg^{2+} . The highest measured calcium concentration is 91 mg/l and the lowest is 30 mg/l, and the mean concentration is 64.92 mg/l. In comparing the

concentration of calcium in the rainy season to that of irrigation season, it is found that the concentration of calcium increased by about 50% in the irrigation season. In the case of magnesium the concentrations during the rainy season range from 6 to 42 mg/l with an average value of 20.15 mg/l. While in dry season the concentrations of magnesium (Mg^{2+}) range from 8 to 41 mg/l with an average value of 22.25 mg/l. In general, without exception Ca^{2+} is dominant over Mg^{2+} and magnesium do not show any significant change in concentration.

During rainy season Na^+ values range from 8 to 140 with a mean concentration of 43.33 mg/l (Fig.5). In dry season the values range from 18 to 193 with a mean concentration of 54.56 mg/l. In the irrigation (dry) season, the concentration of sodium is increased by about 25% compared to the rainy season. K^+ values very low and do not show any variation and hence not considered here.

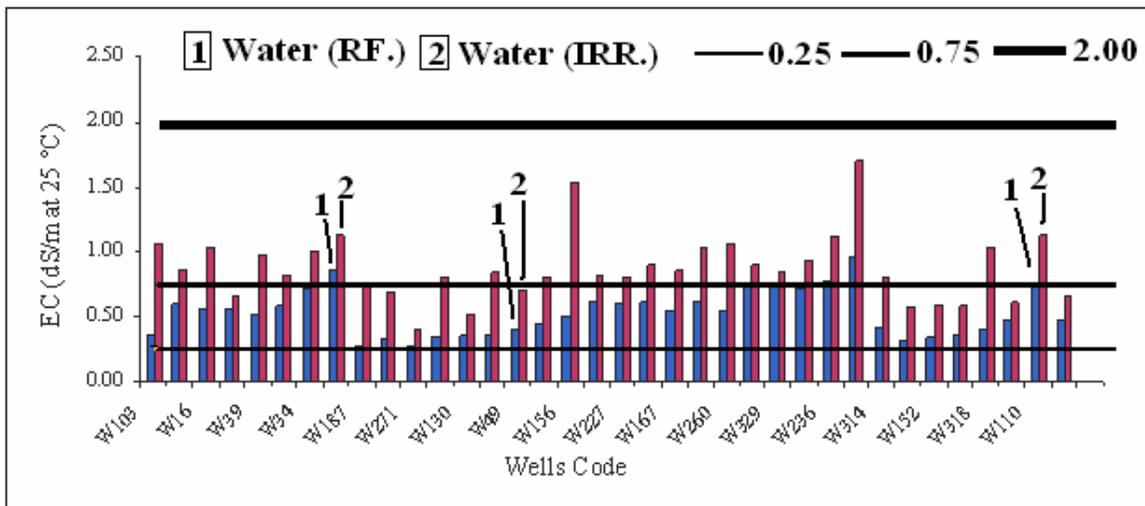


Figure 4. EC of groundwater samples in rainfall and irrigation season (RF- rainy season; IRR- Irrigation (dry) season).

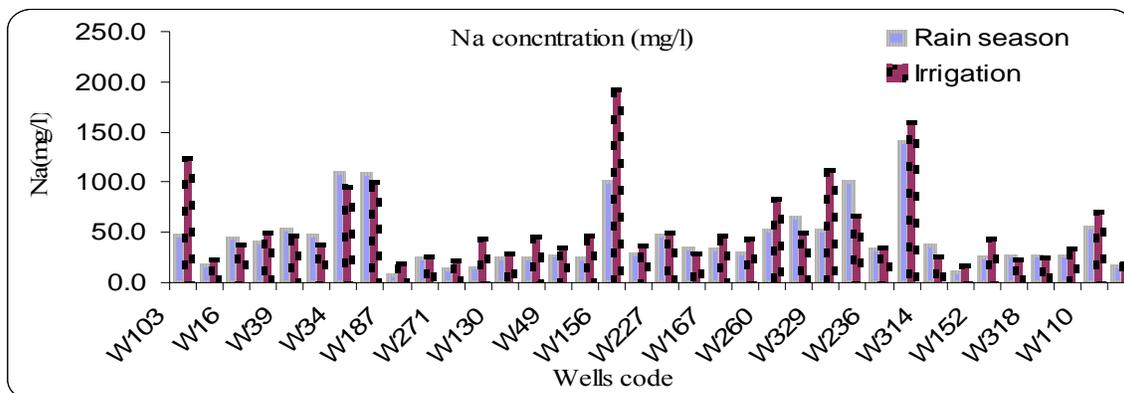


Figure 5. Na concentration in the rainy and irrigation seasons.

4.3 Anions

4.3.1 Bicarbonate (HCO_3^-) and Carbonate (CO_3^{2-})

In the study area, carbonate is found mainly in the form of bicarbonates and is the most abundant ion among anions. During rainy season, the concentrations of bicarbonate (HCO_3^-) range from 82.88 to 477.50 mg/l with an average value of 230.28 mg/l. The same during dry season range from 170.80 to 547.00 mg/l with a mean value of 342.38 mg/l. The mean concentration of bicarbonate in irrigation season is increased by about 12% compared to the rainy season. In both seasons the carbonate (CO_3^{2-}) concentrations are very low and all are below 6.7 mg/l.

4.3.2 Chloride (Cl^-)

In the study area chloride is the second abundant anion next to bicarbonate. Chloride values range from 9.5 to 133.4 mg/l in the rainy season with a mean concentration of 50.71 mg/l. In the irrigation season, the concentrations of chloride (Cl^-) range from 6.5 to 156.1 mg/l with an average value of 45.41 mg/l. Chloride ion do not show any significant seasonal variation (Fig. 6).

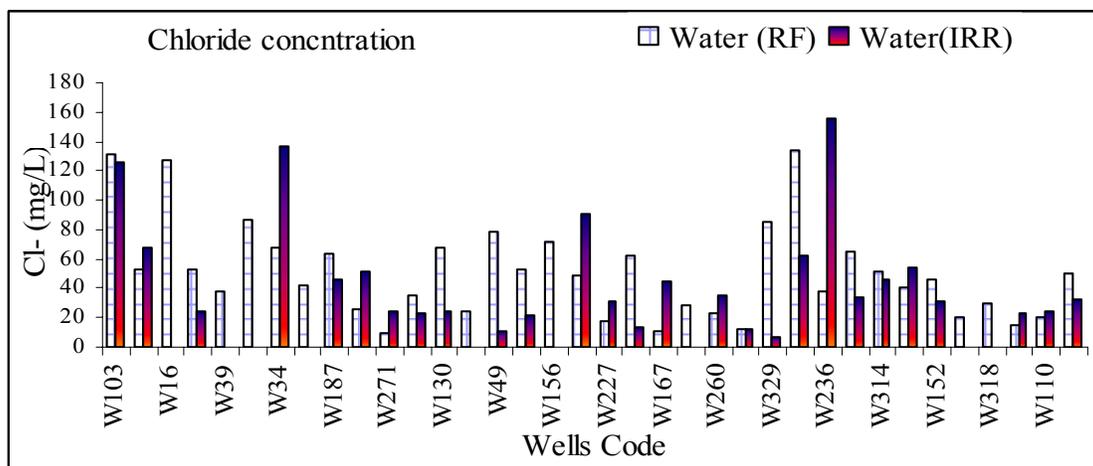


Figure 6. Chloride ion concentration in the rainy and irrigation seasons (RF- Rainy season; IRR- Irrigation (dry) Season)).

4.3.3 Sulphate (SO_4^{2-})

Sulphate is the third most abundant anion next to bicarbonate and chloride in the study area. The values range from 7.2 to 51.80 mg/l in the rainy season with a mean concentration of 25.03 mg/l. In the irrigation season, the concentrations of sulphate (SO_4^{2-}) range from 15.12 to 86.09 mg/l with an average value of 36.35 mg/l.

4.3.4 Nitrate (NO_3^-)

The measured nitrate concentrations range from 0.5 to 43.9 mg/l in the rainy season with a mean concentration of 9.74 mg/l and range from 3.00 to 57.98 mg/l for dry season with an average value of 16.47 mg/l (Fig.7).

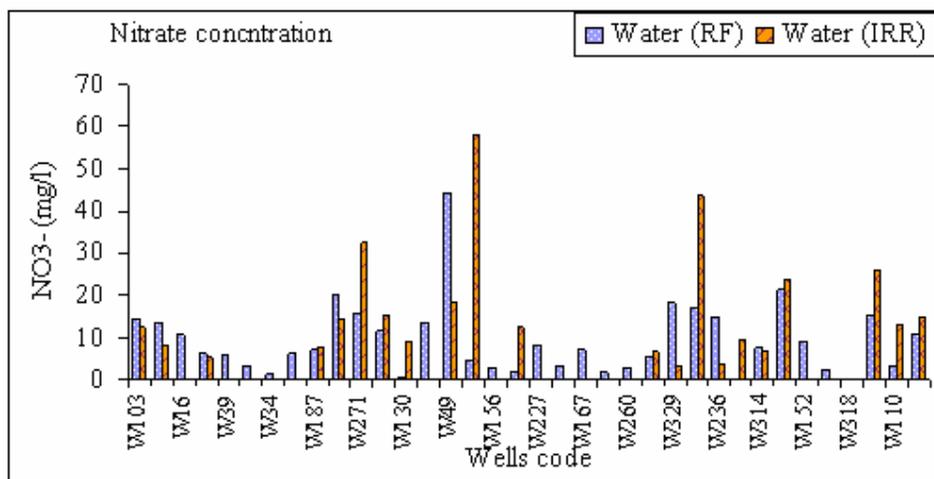


Figure 7. Nitrate values in water from hand dug wells (RF- rainy season; IRR- dry season).

4.4 Water Type

To understand the chemical type of water, different diagrams such as Piper, Stiff, and Box and Whisker diagrams were used to plot the data. The purpose of plotting data in various diagrams is to see how much effective they are for irrigation data and what additional information they provide compared to Piper diagram. Before plotting the proportions of both the anions and cations they were converted from ppm to epm. Samples collected in both rainy and irrigation season were plotted in these Piper, Stiff and Box and Whisker diagrams. They are shown in figure 8 (Piper); figure 9 (Stiff) and figure 10 (Box and Whisker). According to these plots groundwater type in the study area varies from Mg-HCO_3 , Ca-HCO_3 , Mg-Cl , Na-HCO_3 to K-HCO_3 type. During rainy season, the dominant water type is Ca-HCO_3 and is indicated by 42% of the samples, followed by Mg-HCO_3 by 30%, Na-HCO_3 by 22%, Mg-Cl by 3% and K-HCO_3 by 3% samples. In irrigation season the Ca-HCO_3 is the dominant type indicated by 69% of the samples and is followed by Na-HCO_3 type by 24% and Mg-HCO_3 type by 7% samples. Interestingly, Piper diagram shows HCO_3^- and mixed cation type for rainy season and Ca-Na-HCO_3 type for dry season. Whereas Stiff diagram shows Ca-HCO_3 type for rainy season and Na-HCO_3 for dry season and Box and Whisker plot indicates Ca-HCO_3 type for both the seasons.

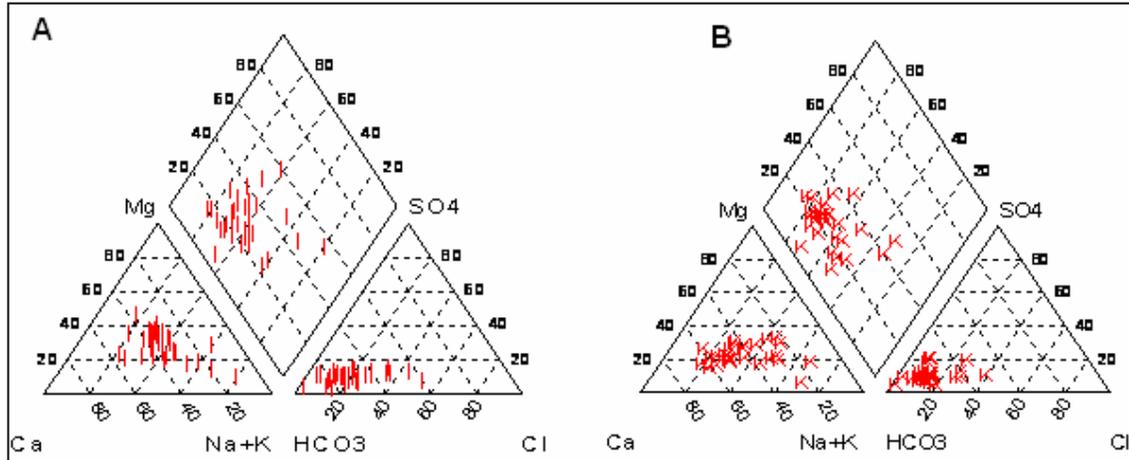


Figure 8. Piper diagram for the water samples A. rainy and B. irrigation season.

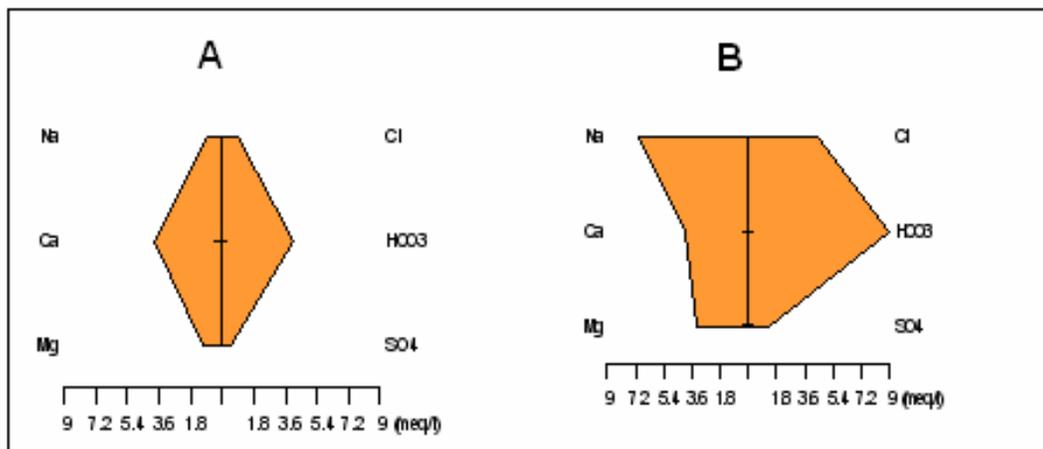


Figure 9. Stiff diagrams (A) Ca -HCO₃ water type and (B) Na-Ca-Mg-HCO₃-Cl type.

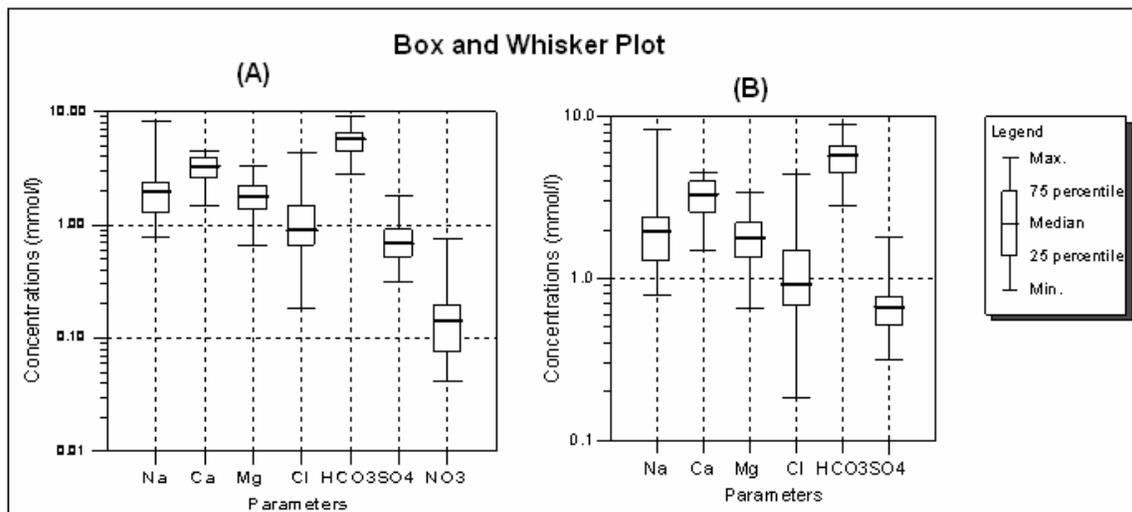


Figure 10. Box and Whisker plot for water samples, (A) rainy and (B) irrigation season.

4.5 Sodium Absorption Ratio (SAR)

SAR (Todd, 1980) was calculated using Aqua Chem 4.0 software. The values range from 0.35 to 6.51 with an average value of 1.51 for rainy season. The mean, minimum and maximum SAR during the irrigation season is 1.44, 0.49 and 5.63, respectively.

5. DISCUSSION

5.1 Physical and Chemical Characteristics

Groundwater being subsurface water derives its contents from different sources like soil, weathered rock, fresh rock and fertilizer with which it will be in contact during infiltration or percolation. This will be aided by the climate (ex. Temperature); CO₂ either free or dissolved state in soil and time duration. Variations in the content of groundwater often associated with irrigation practices, temperature, CO₂ and time as the main lithological source generally remain constant. Such changes are prominent during rainy and dry seasons where both quality and quantity will be affected. During rainy season supply of water being more its ability to carry suspended material will be more compared to ions due to limited amount time duration. The data collected on color, taste etc in the field, shows that about 12% of the samples are muddy or turbid in rainy season with red/brownish color due to the presence excessive amount of suspended matter. Whereas it was found that the water was transparent in dry season. Regarding other parameters, there was no problem of odor or taste or aroma in samples in both the seasons.

Furthermore, a major part of the water in rainy season leaves the catchment as run off and the remaining part will percolate through irrigated soil, alluvial cover and weathered rock to recharge groundwater. So the shallow hand dug wells get recharge prominently from two different routes, 1) drainage channels with sand dominated channel sediments, and 2) irrigated soil, non-irrigated soil and weathered rock. Thus sheet flow and stream water supply dominate rainy season while soil infiltration during dry season. This is well reflected in the compositional variations observed in the water samples collected during rainy and dry seasons. Such as Ca, HCO₃, NO₃ and SO₄ are dominant in the dry compared to the rainy season. The reason for such variations is partly natural and partly anthropogenic.

EC, pH and TDS data suggest though shows variation in their values the majority of the values lie within the acceptable range for irrigation purpose (FAO, 1989). According to FAO (1989)

irrigation water quality guidelines, normal pH range is 6.5 to 8.4. Out of the total analyzed groundwater samples in both sampling seasons, 97% of them are within this normal range.

In the case of carbonate and bicarbonate ions, which are derived from the dissolution of carbonate minerals and available carbonic acid in the soil horizons, are within the acceptable range according to FAO (1985) water quality guidelines for irrigation purpose. The acceptable range for bicarbonate acceptable range is 1-10 meq/l for bicarbonate and 0-0.1 meq/l for carbonate. Therefore, the groundwater of the watershed has no restriction in use for irrigation purposes.

Chloride values being <4 meq/l in the samples collected in rainy season and also the same values in all the samples except one for dry season indicating that if the groundwater of the watershed is considered for surface irrigation use, no chloride toxicity will rise. However, a groundwater from one well needs slight to moderate restriction on use.

According to FAO (1985), water with a sulphate concentration of 0-20 meq/l is considered as a usual range in irrigation water. In the study area, the concentration is not above 1.0 meq/liter. Therefore, it is recommended for use for irrigation purpose without restriction.

Unlike most other elements in groundwater, nitrate is not derived primarily from the minerals (rocks) that make up the groundwater reservoir. Instead, nitrate enters groundwater as part of nitrogen cycle in the earth's hydrosphere and biosphere. In groundwater natural nitrate concentrations generally range from 0.1 to 10 mg/l (Davis and Dewiest, 1966). Nitrate - nitrogen ($\text{NO}_3 - \text{N}$) content greater than 30 mg/l is considered harmful and is severely restricted to use for irrigation purposes (FAO, 1989). Out of the total analyzed groundwater samples (Fig. 8), 31% in rainy season and 3% in irrigation season have a concentration of nitrate - nitrogen ($\text{NO}_3 - \text{N}$) below 5 mg/l. Groundwater from these wells is not hazardous and needs no restriction on use (FAO, 1989). 3% in rainy season and 8% in irrigation season have a concentration of nitrate - nitrogen ($\text{NO}_3 - \text{N}$) greater than 30 mg/l. Groundwater from these wells is hazardous and needs severe restriction on use. In the remaining 66% in rainy season and 89 % in irrigation season, the groundwater of the watershed have a concentration of nitrate - nitrogen ($\text{NO}_3 - \text{N}$) between 5 and 30 mg/l. Groundwater from these wells needs slight to moderate degree of restriction on use. The source for this ion seems to be anthropogenic.

In the case of Ca, Mg and Na the concentration being in the range of 0-20 meq/l for Ca^{2+} , 0-5 meq/l for Mg^{2+} and 18-193 ppm for Na. These values though are within the acceptable range for

irrigation (FAO, 1985) purpose, some of the samples showing higher values for Na in dry season which is well reflected in the Stiff diagram (Fig. 9), seems is an indicator towards increasing Na with time and in turn SAR. Sodium ion is mainly derived from the breakdown of silicate minerals particularly feldspars (albite) and generally get mobilized in water in dissolved state and hence will be lower its content in the stream and shallow hand dug well as well in the rainy season. It is also well reflected in its concentrations during rainy season. But during irrigation (dry) season the values are higher because the well gets its recharge mainly from the soil and rock sources in addition from drainage channels. The reason for such increase of Na content may be because 1) it gets more time to interact with the source material soil and rock; 2) the irrigated soil is becoming sodium-rich gradually due to its absorption replacing Ca; 3) availability of more minerals in the weathered front and alluvium for interaction and dissolution. Though the reasons are common for other cations as well in the case of Na, absorption process seems to be prominent.

5.2 Salinity Hazard

In the rainy season, out of the analyzed thirty six groundwater samples, twenty eight samples have an electrical conductivity values below 0.7 dS/m and the remaining eight samples have an electrical conductivity ranging from 0.72 to 0.96 dS/m (Fig. 4). In irrigation (dry) season, out of the analyzed thirty six groundwater samples, eleven samples have an electrical conductivity values below 0.7 dS/m and the remaining twenty five samples have a range from 0.71 to 1.70 dS/m. Therefore, based on electrical conductivity values, two types of groundwater are recognized in the study area- groundwater which is not hazardous and needs no restriction on use and groundwater which needs slight to moderate degree of restriction on use. The first type groundwater can be used for irrigation for almost all crops and for almost all kinds of soils. No soil or cropping problems will rise. Very little salinity may develop which may require slight leaching; but it is permissible under normal irrigation practices except in soils of extremely low permeabilities. To achieve a full yield potential using the second type, gradually increasing care in selection of crop and management alternatives are required.

5.3 Water Infiltration Rate (Sodicity)

An infiltration problem related to water quality occurs when the normal infiltration rate for the applied water or rainfall is appreciably reduced and water remains on the soil surface too long or infiltrates too slowly to supply the crop with sufficient water to maintain acceptable yields.

Although the infiltration rate of water into soil varies widely and can be greatly influenced by the quality of the irrigation water, soil factors such as structure, degree of compaction, organic matter content and chemical make-up can also greatly influence the intake rate. The two most common water quality factors which influence the normal infiltration rate are the salinity of the water and its sodium content relative to the calcium and magnesium content. High salinity water will increase infiltration. Low salinity water or water with high sodium to calcium and magnesium ratio will decrease infiltration. Both factors may operate at the same time. The infiltration rate generally increases with increasing salinity and decreases with either decreasing salinity or increasing sodium content relative to calcium and magnesium - the sodium adsorption ratio (SAR). Therefore, the two factors, salinity and SAR provide information on its ultimate effect on water infiltration rate.

5.3.1 Rainy season

Out of the analyzed thirty six groundwater samples, 1) five samples show SAR values between 3 and 6 and electrical conductivity between 1.2 and 0.3 dS/m at 25 °C; 2) twenty three samples have a computed SAR ranging from 0 to 3 and an electrical conductivity value ranging from 0.7 to 0.2 dS/m at 25 °C; and 3) seven samples have a computed SAR values ranging from 0 - 3 and a corresponding electrical conductivity greater than 0.7 dS/m at 25 °C. In this season in general two types of water are recognized: groundwater which needs slight to moderate degree of restriction on use (i.e., groundwater from both the five and twenty three samples) and a groundwater which is not hazardous and needs no restriction on use, which is a groundwater from the last seven samples (FAO, 1989).

5.3.2 Irrigation season

Out of the analyzed thirty six groundwater samples, 1) eleven samples have SAR values ranging from 3 to 6. However, these eleven samples are classified in to two groups based on their respective electrical conductivity values. Of the eleven samples two of them have an electrical conductivity greater than 1.2 dSm⁻¹ at 25 °C, indicating no restriction on use and in the nine samples their respective electrical conductivity values are in the range of 1.2-0.3 dSm⁻¹ at 25 °C, indicating that the groundwater from these hand dug wells needs slight to moderate restriction on use; 2) the remaining twenty five samples have a computed SAR values ranging from 0 to 3. However, these twenty five samples are classified in to two groups based on their respective electrical conductivity values. Out of the twenty five samples, eighteen samples have an

electrical conductivity value greater than 0.7 dSm^{-1} at $25 \text{ }^\circ\text{C}$, indicating no restriction in using this water, and the remaining seven samples have an electrical conductivity value ranging from 0.7 to 0.2 dSm^{-1} at $25 \text{ }^\circ\text{C}$, indicating that the water needs slight to moderate degree of restriction on use.

In addition, Wilcox diagram (Wilcox, 1955) was also used to evaluate the potential of the groundwater to create hazard of salinity and sodicity. Accordingly, three classes of water are recognized in the rainy season: S_1-C_1 (low Sodium/Alkali and low salinity hazards), S_1-C_2 (low Sodium/Alkali and medium salinity hazards) and S_1-C_3 (low Sodium /Alkali and high salinity hazards). As shown in figure 11, out of the analyzed samples 14% of them are in the class of S_1-C_1 (low Sodium /Alkali and low salinity hazards), 80% in the class of S_1-C_2 (low Sodium /Alkali and medium salinity hazards) and the remaining 6% in S_1-C_3 (low Sodium /Alkali and high salinity hazards).

In the irrigation season two classes of water are recognized: S_1-C_2 and S_1-C_3 . As shown in figure 12, out of the analyzed samples 30 % of them are in the class of S_1-C_2 and the remaining 70 % are in the class of S_1-C_3 .

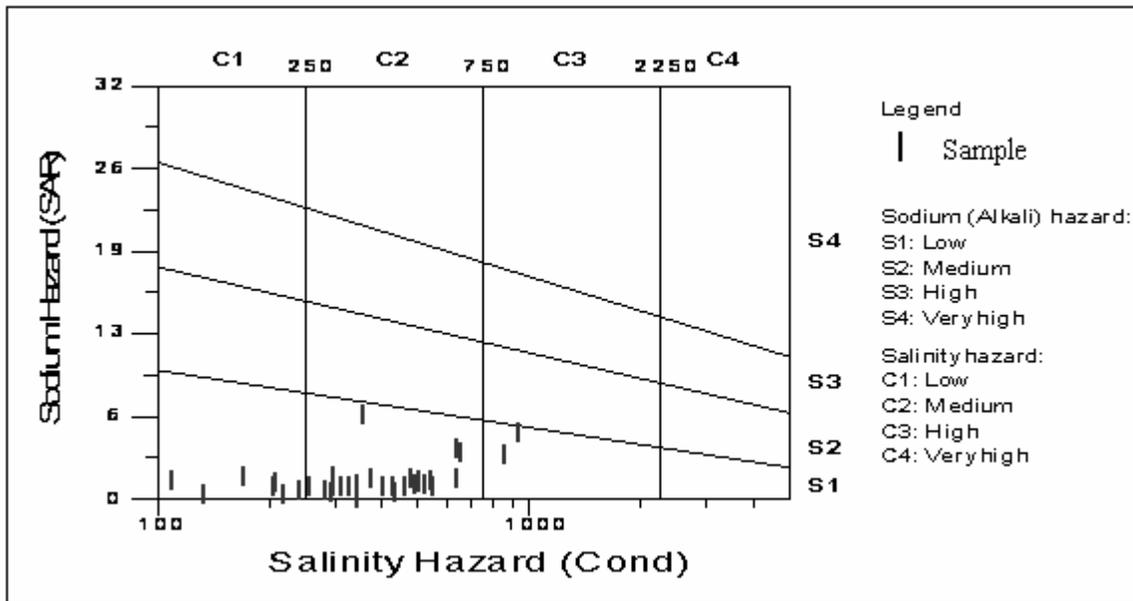


Figure 11. Irrigation water classes using Wilcox diagram (Rainy season).

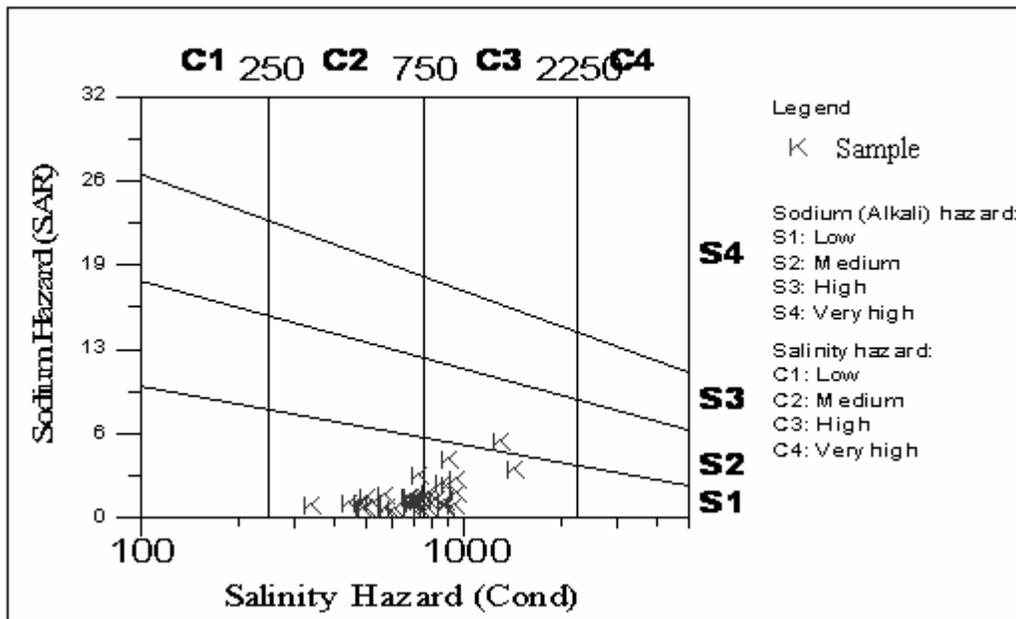


Figure 12. Irrigation water classes using Wilcox diagram (Irrigation season).

5.4 Toxicity Problems

5.4.1 Rainy season

With the exception of five samples, in all the remaining thirty one samples the SAR values are well below 3, indicating no sodium toxicity will rise by using the groundwater from these hand dug wells for surface irrigation (FAO, 1989). In the five samples the SAR values lie within the range from 3 to 9, suggesting the necessity of slight to moderate degree of restriction on use of the groundwater from these five hand dug wells for surface irrigation. The likelihood of sodium toxicity hazards is high if the groundwater from these five hand dug wells is considered for use.

In the case of SAR, the values below 9 indicate low or little danger for sodium and use on sodium sensitive crops must be cautioned; between 10 and 17 medium hazard; 18 and 25 highly hazard and above ≥ 26 very high hazard (Grattan, 2002).

5.4.2 Irrigation season

Out of the analyzed thirty six groundwater samples, in the twenty five samples the SAR values are well below 3, indicating no sodium toxicity will rise by using the groundwater from these hand dug wells for surface irrigation (FAO, 1989). In the remaining eleven samples the computed SAR values lay within the range from 3 to 9, suggesting the necessity of slight to moderate degree of restriction on use of the groundwater from these eleven hand dug wells for

surface irrigation. The likelihood of sodium toxicity hazards is high if the groundwater from these eleven hand dug wells is considered for use.

5.5 Water Classes

The suitability of the groundwater for irrigation is also examined based on Wilcox (1955) classification scheme. Accordingly, in the rainy season the groundwater is categorized into two groups: good and permissible. Out of the total analyzed samples, 89% of the groundwater is in good class and the remaining 11% is in permissible class. In the irrigation season, out of the total analyzed samples, 30% of them are in good class and 70% is in the permissible class.

Table 1. Quality classification of water for irrigation (Irrigation season) (after Wilcox, 1955).

<i>EC</i> (<i>dS/m</i>)	<i>Mean</i>	<i>STDEV</i>	<i>No.</i> <i>Wells</i>	<i>Water Class</i> (<i>Todd, 1980</i>)	<i>Degree of restriction</i> (<i>FAO, 1985</i>)	<i>%</i>
0.25-0.75	0.59	0.11	11	Class 2, Good	None	30%
0.75-2.00	0.99	0.22	25	Class 3, Permissible	Slight - moderate	70%

The data generated in this study are limited. The effect of lithologies on the chemistry of water and also seasonal variation could not be understood because of non-exposure of the rock and the wells studied are mainly confined to very shallow hand dug wells not deeper wells. So the data is discussed mainly in terms of its suitability for irrigation purposes and possible toxic effects if any.

6. CONCLUSIONS

In general, in most part of the study area the groundwater quality is suitable for irrigation. If the groundwater is considered for irrigation use, no serious soil or cropping problems will arise. In few areas, however, the groundwater needs slight to moderate degree of restriction on use. In these areas farmers has to be advised on how they should manage and utilize the groundwater for their irrigation activities. Gradual increase of Na due to absorption may cause sodicity problem hence necessary measures are to be taken to control it.

To maximize the opportunity of the utilization hand dug wells for irrigation, the following recommendations are suggested:

- Detail groundwater potential of the watershed should be investigated to overcome the overexploitation of groundwater and to know the maximum limit of abstraction;
- Soil quality analysis should be conducted to see the effect of groundwater utilization for irrigation on soil quality;
- Recharging measures should be given due attention by the community to maintain and maximize the discharge; and,
- Impact assessment study is suggested to see the significance of hand dug wells utilization on livelihoods of the households.

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